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Numerical modelling - hazard and uncertainty

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Modelling of landslide tsunamis



Landslide dynamics



Tsunami generation



Tsunami propagation and inundation





Rotational slumps: Short time scales



Landslide tsunami generation mechanisms Great variety → Dynamics of landslides of major importance for tsunami-genesis

Staged, retrogressive motion: Slow onset, less efficient



Submarine debris flows: Large volumes



Subaerial landslide tsunamis

- Characterised by violent impact, cratering, and splashing
- More local than earthquake tsunamis
- **7** Strong influence of:
 - Frontal slide area (and volume)
 - Landslide velocity during impact
- Landslide dynamics important for the tsunami generation





Subaerial landslide tsunamis – key examples

18th century Japan volcano flank collapses

Oshima-Oshima 1741

Volume 2.5 km³

~2000 fatalities

Shimabara Bay 1792

Volume ~0.5 km³

>4000 fatalities

Most fatal landslide tsunami in history



Western Norway: Loen 1905, 1936 and Tafjord 1934 174 Fatalities



Many recent high run-up landslide tsunami events: Paatuut, Greenland, 2000 Stromboli, 2002 Aysen fjord, Chile, 2007 Chechalis Lake, Canada, 2007 Askja, Iceland, 2014 Taan fjord, Alaska, 2015 Yangtze River, 2015 Karrat fjord, Greenland, 2017 Anak Krakatau, 2018



Lituya Bay 1958 > 500 m run-up

Highest tsunami run-up height globally





Norway Tafjord 7. april 1934





Church boat MB Tafjord





INGVALD MØLLERSTAD/AFTENPOSTEN/SCANPIX; NTB/Scanpix; http://www.dagbladet.no/magasinet/2008/03/27/530700.html

GIS method for hazard evaluation

Norway:

- 200.000 lakes/reservoirs + 25 000 km coastline
 - 20.000 lakes > 0.1 km²
- Available data
 - Topography and maps of all lakes
 - Landslide data (limited)



Mountain side north of lake Zakarias, Norddal. Potential rockslides identified by NGI (2004)

GIS rockslide tsunami hazard analysis

- **GIS** method for mapping potential rock slide / hazard in fjords, lakes and reservoirs
- Topography steep enough? Rockslide sources large enough?
- Run-out ratio H/L vs. volume
- Probability of certain volumes
- → Topographic rock slide potential for each lake
- **What lakes should be analysed further**
- Co-sponsored by NVE





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Romstad, Harbitz & Domaas (2009) NHESS 9

Results

- ~12 000 lakes (6%) have a tsunamigenic landslide potential
- Most lakes provide a low score (< 0.1% of the maximum)
 - 100 lakes with score >10% of maximum
 - 46 of these are hydro electric power dams;
 - 20 associated with possible large consequences
- Lakes with known hazard (e.g. Lovatnet) are among the lakes with very high score



Landslide dynamics models

- Block models
- Cohesive fluid dynamics models (clay dominated)
 - E.g. Bingham fluids, Hershel-Bulkley models
- Frictional collisional fluid dynamics models (particle interaction dominated)
 - Wide range of models, rheologies, and complexity (e.g. Savage-Hutter, μ(I), pore fluid effects, entrainment, 2D, 3D...)
- Important for submarine landslides
 - Hydrodynamic resistance





Saharan sand dominated slides, Masson et al. (2006)



Tsunami models

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- Shallow water type models 2HD
 - Efficient and most used, but lack essential aspects important for landslide tsunamis
- Boussinesq type models 2HD
 - Include frequency dispersion
- Computational Fluid Dynamics (CFD) models – 3D
 - Three dimensional with few simplifying assumptions
 - Landslide dynamics, complex rheology, and tsunami generation can be fully integrated





CFD model in OpenFOAM

- As PhD student at NGI, Matthias Rauter developed a novel landslide tsunami model (funded by H2020 EU project SLATE)
- Models both the landslide and wave (generation, propagation and runup including water/particle interaction)
- Basis for filling gap in basic physical understanding



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Main aspects of the model Rauter et al. (2022)

Multi-phase coupling, porous landslide



Advanced landslide rheology from solid to granular behaviour

...

$$\begin{split} \nu_{\mathrm{g}} &= \mu(I) \, \frac{p_{\mathrm{s}}}{2 \, \overline{\phi} \rho_{\mathrm{g}}} \frac{1}{\|\mathbf{S}_{\mathrm{g}}\|}, \qquad \mu(I) = \mu_{\mathrm{s}} + \frac{\mu_{\mathrm{d}} - \mu_{\mathrm{s}}}{I_0/I + 1}, \\ I &= \frac{2 \, d \, \|\mathbf{S}_{\mathrm{g}}\|}{\sqrt{p_{\mathrm{s}}/\rho_{\mathrm{g}}}}. \qquad \mathbf{S}_{\mathrm{g}} = \frac{1}{2} (\nabla \mathbf{u}_{\mathrm{g}} + (\nabla \mathbf{u}_{\mathrm{g}})^T) - \frac{1}{3} \nabla \cdot \mathbf{u}_{\mathrm{g}} \mathbf{I}, \end{split}$$

Simulating lab scale experiments



To full 3D simulations



Main scientific findings

- Matching consistently both landslide AND tsunami observations from the laboratory to the field scale
- Close agreement with both landslide run-out and wave observations
- Advanced landslide material behaviour, direct simulation with no attempt to calibrate the landslide parameters







The Åknes rockslide

- Unstable rock slope 150 900 m.a.s.l
- Large movements /deformations
- Largest volume > 50 Mm³
- Advanced computational tools needed
- Laboratory experiments 2D and 3D
 - Calibration and verification of numerical models
- A large number of scenarios and locations analysed



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Åknes tsunami – run-up in Hellesylt



Åknes – laboratory experiments

- 2D: Hydrodynamic laboratory, University
 of Oslo
- 3D: SINTEF Coast and harbor laboratory
 - Scale 1:500
 - "block-slide"

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 Validation of and input for numerical models





Laboratory experiments vs. numerical models

- Laboratory model inside red box
- Measured surface time histories both in fjord and in inundation area
- Demonstrate governing parameters for propagation and model benchmarking
- No tuning of numerical model

 reproduce laboratory
 experiments as close as
 possible







Ξ

elevation

surface

Dam overtopping

- laboratory experiments vs. numerical modelling

- Master thesis
 - Ragnhild Hammeren (2016), NTNU, Trondheim
- Rebuild of Åknes model
- Numerical modelling by NGI





Testing landslide tsunami models

http://www1.udel.edu/kirby/landslide/problems.html

- Benchmark cases (NTHMP US)
 - Mainly idealized laboratory
 - Tailored towards tsunami-genesis, less on landslide dynamics and material behavior
- Friction of landslides depends on its size, need modelling well documented full scale events
 - Lake Askja 2014: confined in lake, accurate volume and runup measurements, "full scale lab"
 - Other examples, Taan fjord, Anak Krakatoa, ...









0.25

0.15

0.25

0.20 ·





Landslide Probabilistic Tsunami Hazard Assessment (LPTHA)

30

29.5

Latitude (degrees)

27.5

27

LPTHA in brief:

- Generate a synthetic set of sources (different volumes and generation mechanisms)
- Define annual source probabilities – use statistics of past data
- 3. Simulate the wave propagation for each source
- 4. Aggregate probabilities from all simulations to hazard curves



Landslide Probabilistic Tsunami Hazard Analysis (LPTHA) example

Løvholt et al., 2020, Landslides

- Goal estimate tsunami probability of occurrence
- Include the uncertainties in forecasting
- Area of interest Lyngen Norway
- Four different unstable rock slopes
 - Volumes 0.8-6 Mm³
 - Frequencies estimated prior to our analysis by means expert judgement
 - Average frequencies 1/633 yr⁻¹ 1/2315 yr⁻¹



Schematic LPTHA



Establishing and aggregating uncertainties

Løvholt et al., 2020, Landslides

- Epistemic (systematic) uncertainty landslide dynamics
 - Block slide much more controlled behaviour of the slide (shape, run-out, velocity)
 - Run-out distance R fitted towards past run-out
 - For the impact velocity U and frontal areas A: Sensitivity studies based on modelling and experience from modelling past events
- Event tree analysis

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Rates from all possible landslide events

Uncertainty in landslide run-out distance (H/L) from past data





LPTHA results for Lyngen – local inundation analysis aggregated: 31 locations x 600 events

Løvholt et al., 2020, Landslides

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Example Stiksmoen, Flåm – western Norway

- **4**00.000 m³
- Estimated annual probability of release, p=1/240
- Maximum line highest runup of all 600 scenarios (p -> 0)
 - Worst case scenario?
 - Used for evaluation of evacuation zones and locations for critical infrastructure (e.g. hospital, p=0)



Calibration of LPTHA

- LPTHA for Norwegian landslide tsunamis
- Ongoing work
- Using known and documented events
- **7** Systematic analysis
 - Common correction factors to frontal area?
 - Calibrate the LPTHA so that the event i close to the median values of the runup?
- Using both the OpenFOAM model in addition to depth averaged models for landslide and tsunami



- ✓ Skafjell, 1731
- ✓ Tjelle, 1756
- ✓ Taford, 1934
- ✓ Rissa 1978
- ✓ Årdalstangen, 1983
- ✓ Katlenova, 1998
- ✓ Statland, 2014
- ✓ Geiranger, 2017

Tsunamis in reservoirs

- Main modelling issues in a fjord are also present in a lake/reservoir
- **7** Tsunami origin both from subaerial and submarine landslides
- Models for tsunami run-up can handle dam overtopping
 - Caclulate forces on dam itself and infrastructures on dam crown during overtopping
 - Can also be coupled to modelling of the downstream flooding
- LPTHA is a good methodology for treating uncertainties systematically also for reservoirs
 - Most likely or worst case scenario, or a given return period?

Conclusions

- Landslide dynamics accounts for large variability
 - Wide range of sources and mechanisms more complex and diverse than earthquakes
- Systematic uncertainty is linked to material parameters and hydrodynamic resistance
- Hindcasting past landslide run-out implies large uncertainties
- Tsunami data can be used to narrow down this uncertainty range
- With better models, we are on the pathway to understand well the generation mechanisms, but difficulty remains
 - Scaling up from the lab
 - Incorporating real geomaterial behaviour and phase transitions
- LPTHA is the most rational tool to manage these uncertainties in forecasting



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